Stress orientations of Taiwan Chelungpu-Fault Drilling Project (TCDP) hole-A as observed from geophysical logs

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1. Introduction

The Taiwan Chelungpu-fault Drilling Project (TCDP) drilled a 2-km-deep research borehole to investigate the structure and mechanics of the Chelungpu Fault that ruptured in the 1999 Mw 7.6 Chi-Chi earthquake. Geophysical logs of the TCDP were carried out over depths of 500–1900 m, including Dipole Sonic Imager (DSI) logs and Formation Micro Imager (FMI) logs in order to identify bedding planes, fractures and shear zones. From the continuous core obtained from the borehole, a shear zone at a depth of 1110 meters is interpreted to be the Chelungpu fault, located within the Chinshui Shale, which extends from 1013 to 1300 meters depth. Stress-induced borehole breakouts were observed over nearly the entire length of the wellbore. These data show an overall stress direction (~N115°E) that is essentially parallel to the regional stress field and parallel to the convergence direction of the Philippine Sea plate with respect to the Eurasian plate. Variability in the average stress direction is seen at various depths. In particular there is a major stress orientation anomaly in the vicinity of the Chelungpu fault. Abrupt stress rotations at depths of 1000 m and 1310 m are close to the Chinshui Shale’s upper and lower boundaries, suggesting the possibility that bedding plane slip occurred during the Chi-Chi earthquake. Citation: Wu, H.-Y., K.-F. Ma, M. Zoback, N. Boness, H. Ito, J.-H. Hung, and S. Hickman (2007), Stress orientations of Taiwan Chelungpu-Fault Drilling Project (TCDP) hole-A as observed from geophysical logs, Geophys. Res. Lett., 34, L01303, doi:10.1029/2006GL028050.

2. Borehole Breakout and Fracture Orientation From Image Logs

Borehole breakouts result from localized failure around a borehole in response to horizontal compression [Bell and Gough, 1979; Zoback et al., 1985]. In a vertical borehole, the breakout directions are parallel to the minimum horizontal principal stress (Shmin). Due to the symmetry of the stress concentration around the borehole, breakouts occur in pairs, offset by 180°. Figure 2 presents an example formation micro image from FMI for the depth of 1408.9–1411.3 m. The breakouts appear as dark zones of lower electrical conductivity 180° apart. Localized breakout rotations indicate stress perturbations associated with slip on active fault intersected by the borehole [Barton and Zoback, 1994]. Thus, by investigating localized changes in breakout direction, active fault zones can be identified. Figure 3a shows a profile of the orientation of the minimum horizontal principal stress (Shmin) from 500–1860 m (solid cycles) based on wellbore breakout directions averaged over 2 m depth intervals. The borehole azimuth (open cycles) during drilling is shown for reference. As shown in Figure 3c, the borehole deviation from vertical is very small (<3°) except near bottom of the hole where it deviated slightly. Comparing the breakout direction and hole azimuth can help us to determine whether the failures result from changes in stress.
Figure 1

(a) Map showing the location of TCDP (Dakeng) and the surrounding geology. The map includes the Chelungpu Fault and Chichi fault zones.

(b) Cross-section diagram of the Chelungpu Fault and TCDP (Dakeng) area, highlighting the stress orientations and geological layers.
direction or from drilling-induced erosion of the wellbore wall [see Plumb and Hickman, 1985]. The borehole azimuth differs from the observed orientation of breakout pairs, suggesting that the breakout was resulted from stress rather than drilling. The borehole breakout pairs obtained from the FMS data indicate an average orientation of N25°E over most of the 500–1800 m interval (Figure 3a). This orientation corresponds to a direction of $S_{\text{Hmax}}$ of N115°E (Figure 3b) which is essentially parallel to the regional stress direction in the Taiwan region that results from the convergence of the Philippine Sea Plate to the Eurasian plate [Kao and Angelier, 2001] (see inset of Figure 1a). The deviation of TCDP-A was $< 4^\circ$ from the surface to a depth 1600 m, and only in the lower part of the hole (deeper than 1835 m) did the deviation of the borehole exceed $> 12^\circ$ where almost no breakouts were observed. Hence, the effect of borehole deviation was ignored in this study [Peska and Zoback, 1995].

As shown in Figure 3b, minor fluctuations of the orientation of $S_{\text{Hmax}}$ are seen at a variety of depths, with a major change of the $S_{\text{Hmax}}$ direction seen between 950 m and 1310 m. Below 1500 m, very few breakouts are observed and it is difficult to assess whether the variation of orientations observed are statistically significant.

As mentioned above, the locations of abrupt breakout rotations indicate the presence of active shear zones. In total, five zones with significant rotation of the $S_{\text{Hmax}}$ at the depths of 1000 m, 1110 m, 1230 m, 1310 m, and 1460 m were identified within the depths of 600–1500 m. These zones are concentrated in the depth interval 980–1300 m, corresponding to the lithology in the Chinshui Shale 1500 m (Figure 3). The abrupt stress rotations at depths of 1000 m and 1310 m are close to the Chinshui Shale’s upper and lower boundaries, suggesting the possibility that bedding plane slip occurred during the Chi-Chi earthquake. In addition to these rotations, the most significant changes in the $S_{\text{Hmax}}$ direction occur near depths of 1110 m and 1230 m.

Numerous fractures, faults and bedding-planes are obvious in the image logs. Figure 3c shows the distribution of number of visible fractures and faults in 10-m intervals. 63% of the fractures (4,473 of 7,093) are found between 500 and 1100 m with nearly 50% of the fractures concentrated between 900 and 1100 m. The number of fractures decreases gradually between 1100 m and 1500 m and then more rapidly below 1500 m where they are nearly absent.

### 3. Physical Properties From Geophysical Logs

Geophysical logs were run from 500 m to 1860 m to obtain the physical properties of fault zones and adjacent damage zones, including P & S wave velocities, density, resistivity, and gamma ray. The logs sample these properties every 15 cm. The velocity logs show a gradual increase in P and S wave velocity (Figures 4b and 4c) due to the increase in confining pressure with depth. The P wave (compressional) velocity ranges from 2.27 km/s at depth 551 m to 5.47 km/s at depth 1668 m. The S wave (shear) increases from 1.4 km/s at a depth of 597 m to 2.98 km/s at a depth of 1705 m. Overall average of P-wave velocity is 3.81 km/s and S-wave velocity value is 1.86 km/s. The overall resistivity of the formations encountered in TCDP is low from 20 ohm-m to 50 ohm-m between depths of 587 m to 1856 m (Figure 4d). Depths of abnormal geophysical properties can be identified at depths where the velocities and resistivities are lower than the average trend with depth. The gray bars in the lithologic column represent fault zones observed in the core (E.-C. Yeh, personal communication, 2006). Surface rocks near the drill site are identified as Cholan Formation and Chinshui Shale. The drill site is in the town of Dakeng, 2.5 kilometer east to the Chelungpu fault rupture (wide black dash line). The nearby faults from west to east are Changhua fault, Shuangtung fault and Shuichangliu fault (black dash lines). The regional tectonic stress of Taiwan is N126°E as shown in inset [Kao and Angelier, 2001]. The slip direction of the Philippine Sea Plate with respect to the Eurasian plate is in this same direction. The convergence rate is 82 mm year$^{-1}$. The location of mainshock (asterisk) and its focal mechanism are also shown. (b) A geological cross-section near the drill site [Lyu et al., 2005]. The Chelungpu fault dips 30 degrees east. TCDP drilled a 1.8 km depth hole through from Cholan Formation to Chinshui Shale and Kueichuline Formation, then, to Cholan Formation. The Sanyi fault is interpreted to be at the base of the Kuichulin Formation with some uncertainties shown by dash line below it.
2006). There is some degree of correlation between fault zones and anomalous geophysical properties as observed in other fault zone drilling project [e.g., Boness and Zoback, 2004]. However, this correlation is not compelling as a number of zone of anomalous geophysical properties do not correlate with identified shear zones and some of the shear zones identified in the core are not associated with anomalous properties.

We can also compare the geophysical logs with the zones of localized $S_{\text{Hmax}}$ perturbations seen in Figure 3b, which are shown by shaded in gray bands in Figures 4a–4f. The zone identified from the discontinuous rotation of $S_{\text{Hmax}}$ at the depth of 1110 m corresponds to the most significant changes in Vp, Vs, and resistivity. It is also consistent with the fault zone identified from the core. Considering the observations from all of the various geophysical logs and the faults identified from cores, it suggests that the shear zone at the depth of about 1110 m is the most recent shear zone related to the 1999 Chi-Chi earthquake. Other zones of anomalous breakout rotations could be associated with co-seismic slip on secondary fault planes that moved during the Chi-Chi earthquake or slip that occurred during historical events. The geophysical logs close to the depth of 1680–1707 m show significant changes in Vp, and Vs, which have good correlation to the fracture zone observed from the core. This shear zone could be related to the SanYi fault beneath the Chelungpu fault as shown in Figure 1b. As mentioned above, the abrupt stress rotations at depths of 1000 m and 1310 m are close to the Chinsui Shale’s upper and lower boundaries, suggesting the possibility that bedding plane slip occurred during the Chi-Chi earthquake.

4. Conclusions

Continuous core, geophysical logs data and observations of stress-induced borehole breakout were collected from the entire depth of TCDP hole-A. These data clearly present the overall stress direction at the site, which is consistent with the regional stress field as well as anomalous stress rotations associated with the faults that cut through the hole. Active shear zones are identified at depths
of 1110 m and 1680–1707 m related to Chelungpu fault and SanYi fault, respectively. The shear zone at the depth 1110 m shows the most consistent features with fracture zones observed in core, the rotations of $S_{\text{Hmax}}$ and anomalies in velocities and resistivity, suggest that this fault is most likely related to the 1999 Chi-Chi earthquake. Other observations were found for the rotation of the $S_{\text{Hmax}}$ direction indicate other active faults as well as bedding plane slip at the boundaries of different formations.

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References

Figure 4. (a) Fractures density picked from FMS and FMI in 10 meter intervals. (b) P-wave velocity, (c) S-wave velocity, (d) Resistivity, (e) Density, and (f) Gamma ray and lithology from core descriptions. The zones with localized stress rotations are shown on geophysical logs with thick gray lines. The fault zones identified from the on-site core description (E.-C. Yeh, personal communication, 2006) are shown by gray bars in the lithology column.


Yue, L. F., J. Suppe, and J. H. Hung (2005), Structural geology of a classic thrust belt earthquake: The 1999 Chi-Chi earthquake Taiwan (Mw7.6), J. Struct. Geol., 27, 2058–2083.


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