Analysis of Temperature, Pressure and Reservoir Properties, Case Study of Aluto Langano Geothermal Field, Ethiopia

Teka N. GEBRU, Svanhjörg H. HARALDSDÓTTIR, Sæunn HALLDÓRSDÓTTIR, Lúdvík S. GEORGSSON
Ethiopian Electric Power (EEP), Addis Abeba, Ethiopia
e-mail: teka2005@yahoo.com, tekishmain@gmail.com

Keywords: Temperature inversion, Pressure pivot point, heating period, Welltester, Transmissivity, Storativity, Injectivity index, Skin factor.

ABSTRACT
Aluto Langano geothermal field is located in the Ethiopian main rift valley confirmed by temperatures greater than 200°C at a depth of 1000m. Two geothermal prospects namely Butajira and Meteka geothermal fields out of the 24 geothermal fields are identified recently in 2017/2018. Of the 10 deep wells drilled in the area, two of them are the only directionally drilled in 2015 to a measured depth of 1921 m and 1950 m. The remaining 8 wells are all vertical, drilled in the 1980’s. The objective of this report is to analyse the temperature and pressure in the newly drilled wells (LA-9D and LA-10D) in order to estimate the formation temperature of the geothermal system as well as to estimate the reservoir parameters of the field based on well test analysis of the wells, LA-4 (the old well) and LA-10D (new drilled well).

The formation temperatures of the wells were estimated using the Horner method. The ultimate goal of temperature and pressure logging in geothermal investigation is to determine formation temperature and reservoir pressure. Bottom hole formation temperature of wells LA-9D and LA-10D were estimated by extrapolation of a short term heating period using Horner plot method and the evaluation of formation temperature for both wells suggest that the reservoir temperature is in the range of 300-310°C. The injection test was processed using Welltester software, which is developed at Iceland GeoSurvey to estimate reservoir and well parameters. The data from the injection tests were too scarce to use directly in Welltester, so manually adding data where there were gaps were used, followed by interpolation in Welltester, and gradually gave good results with respect to the quality of the data. Storativity and transmissivity were of the order of magnitude 10^-4, which is similar to Icelandic high temperature geothermal wells. The skin factor was generally negative, which indicates that the permeability in the closest surroundings of the well is higher than farther away. The injectivity index is rather low, between 1.3 and 2.2 (l/s)/bar, which can explain partly why the wells have not been good producers.

1. INTRODUCTION
Geothermal utilization in Ethiopia began in 1886, when “EtegeTayitu” the wife of King Menelik II, built a bathhouse for herself and members of the Showa Royal Court in Addis Ababa near the “Filwoha” an Amharic term which means “boiled water”, hot mineral springs (Wikipedia, 2017). The hot springs are still open to the public. In 1969, geothermal surface exploration started with a regional geo-volcanological mapping and hydrothermal manifestation inventory in most of the Ethiopian Rift Valley and Afar Depression that revealed high and low enthalpy geothermal resources Teklemariam et al. (2000). Then, after several years of exploration activities, which included geological, geochemical and geophysical investigation, about eighteen (18) geothermal prospect areas were selected for further feasibility studies through drilling of exploration wells. To date only 6 prospects have been subjected to feasibility studies, namely Aluto Langano, Tendaho, Corbetti, Abaya, Tulu Moye, Dofan and Fentale geothermal fields, based on their strategic locations, i.e. proximity to the existing grid and population density (Ministry of Mines and Energy, 2008). Prefeasibility study for the two new geothermal prospects namely Butajira and Meteka geothermal fields has been conducted as a result of good anomalies for geothermal is observed.

A total of 10 deep exploratory wells have been drilled at the Aluto Langano geothermal field to a maximum depth of 2500m. Information about the wells is shown in Table 1, 8 of the wells (LA-1 to LA-8) are vertical, drilled between 1981 and 1986 (ELC 2016), while two of them, LA-9D and LA-10D, are the first directional wells in Ethiopia, drilled to a depth of 1920 and 1950 m respectively, completed in 2016 WestJEC (2016). Four productive wells (LA-3, LA 4, LA-6 and LA-8) in Aluto Langaano geothermal field have been supplying steam and brine to operate a binary pilot power plant commissioned in 1999, with a capacity of approximately 7.3 MWe.

Consultancy service agreement between Ethiopian Electric Power (EEP) and a consultant (West Japan Engineering Company, WestJEC) has been signed to construct a 5 MWe Wellhead Power Plant utilizing the two directional wells and the construction works are expected to be completed by early 2019. The grant is from the Japanese Government.

The objective of this report is to assess the formation temperature and some reservoir parameters of selected wells from the Aluto Langano geothermal field, using the Horner method and well test analyses software respectively. Data for wells LA-9D and LA-10D are used for formation temperature assessment based on temperature and pressure logs, which were measured during warm up and after a 3-month discharge of the wells. Data for wells LA-4 and LA-10D are used for well test analysis where the Welltester software developed at Iceland GeoSurvey was used to manipulate injection test data measured in the 1980’s and 2015’s.
2. GEOLOGIC AND STRATIGRAPHIC SETTINGS OF ALUTO LANGANO GEOTHERMAL FIELD

The Aluto Langano Geothermal field is located in the Lakes District, Ethiopian main Rift Valley, about 220 km south of Addis Ababa and covers an area of about 100 km$^2$. It lies between lakes Langano and Ziway and rises to about 690 m above the surrounding AdamiTullu Plain which has an elevation of about 1600 m a.s.l. (Figure 1). Volcanic activity at the Aluto volcanic center is entirely of Quaternary age and initiated with a rhyolite dome building phase intervened by explosive pyroclastic pumice eruptions. The recent nature of these volcanic products indicates a heat source which is still hot enough at depth. An extensive cap rock having large lateral coverage exists at Aluto Langano in the form of lake sediments and associated with overlying pyroclastic.

The stratigraphic setting of Aluto Langano geothermal field shows that its thickness increases westwards from 300 m to 1000 m and consists of three main units; pre-Aluto volcanic products aged 1.5 to 3.5 Ma, Aluto Volcanic Products aged 2 to 150 ka and Sedimentary formations ELC (2016).

Table 1: Information about the exploratory wells located in Aluto Langano geothermal field (ELC, 2016)

<table>
<thead>
<tr>
<th>Well</th>
<th>WH Coordinates</th>
<th>WH Elevation</th>
<th>Total Meas. Depth</th>
<th>Total Vertical Depth</th>
<th>Elevation of well bottom</th>
<th>Coordinates of well bottom</th>
<th>9 5/8” Casing Shoe</th>
<th>7” slotted liner</th>
<th>Notes</th>
</tr>
</thead>
<tbody>
<tr>
<td>LA-1</td>
<td>853,308</td>
<td>474,047</td>
<td>1021</td>
<td>1317</td>
<td>1317</td>
<td>1317</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-2</td>
<td>820,501</td>
<td>469,899</td>
<td>1704</td>
<td>1662</td>
<td>1662</td>
<td>1662</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-3</td>
<td>800,723</td>
<td>477,401</td>
<td>1921</td>
<td>2144</td>
<td>2144</td>
<td>-223</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-4</td>
<td>860,839</td>
<td>478,395</td>
<td>1957</td>
<td>2062</td>
<td>2062</td>
<td>-106</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-5</td>
<td>859,430</td>
<td>476,757</td>
<td>2038</td>
<td>1869</td>
<td>1869</td>
<td>169</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-6</td>
<td>861,728</td>
<td>477,546</td>
<td>1983</td>
<td>2203</td>
<td>2203</td>
<td>2203</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Unproductive well</td>
</tr>
<tr>
<td>LA-7</td>
<td>860,532</td>
<td>476,296</td>
<td>1891</td>
<td>2449</td>
<td>2449</td>
<td>-558</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-8</td>
<td>862,190</td>
<td>476,944</td>
<td>1956</td>
<td>2501</td>
<td>2501</td>
<td>-406</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-9D</td>
<td>860,786</td>
<td>477,863</td>
<td>1958</td>
<td>1794</td>
<td>1794</td>
<td>172</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
<tr>
<td>LA-100</td>
<td>862,846</td>
<td>478,807</td>
<td>1951</td>
<td>1816</td>
<td>1816</td>
<td>140</td>
<td>verticl well</td>
<td>verticl well</td>
<td>Stable well</td>
</tr>
</tbody>
</table>

3. TEMPERATURE ASSESSMENT

Stabilized formation temperature is one of the most important parameters to evaluate for a geothermal reservoir condition, when a well is completed (Hyodo and Takasugi, 1995). Its determination from well logs, requires knowledge of the temperature
disturbance produced by circulating drilling mud or fluid (Kutasov and Eppelbaum, 2005). The formation temperature of the wells is estimated using Horner plot method. The Horner plot method is the most popular method to estimate formation temperature from down hole logging temperature data (Dowdle and Cobb, 1975), whereas Kutasov and Eppelbaum (2005), stated Horner method as the widely used in petroleum reservoir engineering and in hydrogeological exploration to process the pressure-build-up test data for wells producing at a constant flow rate.

The validity of the Horner plot is based on Fourier’s heat conduction equation, which describes the change in temperature (T) as a function of time (t) and position/space (x).

\[
\frac{c_p \rho}{\partial T}{\partial t} = k \frac{\partial^2 T}{\partial x^2}
\]  

(1)

where \(c_p\) = Heat capacity of the material (J/m\(^3\)°C), \(\rho\) = Density (kg/m\(^3\)), \(k\) = Thermal conductivity (W/m°C);

The Horner plot method uses the measured temperature at a given depth from several temperature logs taken at different times. Horner time (\(\tau\)) is given by Equation 2 below

\[
\tau = \frac{\Delta t}{\Delta t + t_k}
\]  

(2)

where \(t_k\) = Circulation time before shut-in (hrs), \(\Delta t\) = The time elapsed since the circulation stops (hrs);

The circulation time (\(t_k\)) is an important parameter in the Horner plot method and should therefore be determined accurately. Since drilling (and thus injection of cold fluid) reaches different depths at different times, the circulation time varies with depth. The temperature recovery data is plotted logarithmically with the Horner time and the temperature will gather up as a straight line at infinite time, \(\tau = 1\) or log \(\tau = 0\). Figure 2 shows an example of Horner plot of the evaluation of the formation temperature at 1200 m depth in well LA-4. It is known that the temperature found from the Horner plot is usually lower than the actual formation temperature (Helga Tulinious, Senior Geophysicist, ISOR, personal communication, 30 Aug 2017).

Figure 2: Example of a Horner Plot to estimate formation temperature of well LA-4 at a depth of 1200 m

3.1. Estimation of Formation Temperature and Pressure Control Point of well LA-9D

It is the first directional drilling of a geothermal well in the history of the country. The well is directed to N70°W and has a maximum inclination of 51°. The kick off point (KOP) is started at 700 m. The surface casing (20") is set to a depth of approximately 60 m then the anchor casing (13-3/8") is installed to 210 m depth then a production casing (9-5/8") is set to a depth of 605 m and finally the slotted liner (7") is installed to a depth of 1921 m.

In well LA-9D, seven small feed zones were located at depth of 760m, 830m, 940m, 1000m, 1140m, 1640m, 1800m, while a large feed zone encountered at depth of 1350m. The feed zones were identified using warm-up profiles between 07/05/2015 – 09/09/2015 and circulation loss data.

The stabilized formation temperature was estimated from the data set of temperature recovery test and temperature logging during shut in condition. The estimated formation temperature increases with depth and reaches about 309°C at the well bottom which is in line with the West JEC 2016 report.
The data points used were selected from the temperature logs at the depth of the feed zones, their estimated formation temperature was determined and then plotted in the same Figure 3. The formation temperature of the well was estimated from these points. The formation temperature appears to increase in the production part of the well 300-311°C.

The pressure profile obtained during the warm up period and the initial pressure profile are shown in Figure 4. The pressure pivot point is located at around 1120 m and the pressure is 89 bar at that point with vertical depth measurement.

3.2. Estimation of Formation Temperature and Pressure Control Point of well LA-10D

Drilling of well LA-10D was started on June 25, 2015 and completed on October 2, 2015. The directional drilling kick off point (KOP) was at 450m to a direction of N43°W with a maximum inclination of 27.75°. The surface casing (20") is set to a depth of approximately 60 m then the anchor casing (13-3/8") is installed to 343 m depth then a production casing (9-5/8") is set to a depth of 807 m and finally the slotted liner (7") is installed to a depth of 1951 m.
In well LA-10D six small feed zones were located at 920m, 1280m, 1320m, 1620m, 1760m, while large feed zone at 700m were encountered (circulation loss zone) with rapid increase in temperature. The feed zones were identified from injection test and from temperature recovery measurements that were taken on October 5, 6, and 24 2015 and also from circulation loss data (Figure 5).

The formation temperature was estimated from the data set of temperature recovery test of October 6 (12H &24H) and 24 October. The estimated formation temperature simply increases with depth and reaches about 305°C at the well bottom 1800 m depth. This is also in line with the West JEC 2016 report. To calculate the initial pressure, the estimated formation temperature profile plotted in figure 6 was inserted in to the PREDYP (part of the Icebox software package). The water level was adjusted in the calculations until the calculated profile matched the pressure pivot point. The pressure match was achieved more or less with water levels at 200 m depth. The pressure pivot point can be observed around 1288m vertical depth with 96 bar pressure, this shows the major feed zone of the well (Fig 6).

In well LA-10D six small feed zones were located at 920m, 1280m, 1320m, 1620m, 1760m, while large feed zone at 700m were encountered (circulation loss zone) with rapid increase in temperature. The feed zones were identified from injection test and from temperature recovery measurements that were taken on October 5, 6, and 24 2015 and also from circulation loss data (Figure 5).

3.3 Temperature Cross Section (Southeast-to-Northwest direction)

The estimated formation temperatures of wells LA-4, LA-9D, LA-10D, LA-6 and LA-8 are shown in a cross-section from SE to NW (Figures 7 and 1). The formation temperature of LA-4, LA-6, LA-9D and LA-10D was derived mostly from the latest (January 2017) logging data while LA-8 is based on data from 1980’s. As clearly observed from the cross section, the hot up flow seems to be in the vicinity of wells LA-6, LA-9D and LA-10D, which reaches the warmest part (>300°C), probably from the main up flow zone of the reservoir. Well LA-4 is farther away from the hotter part of the reservoir, with temperature ranging from 210°C to 230°C while well LA-8 is closer to the hotter part of the reservoir. This can probably be caused by cold water inflow/outflow towards wells LA-4 from eastern escarpment of the rift. The above profile supports the idea, of Kebede et al., (2002), who wrote about the up flow zone location in the vicinity of wells LA-6 and LA-3. The upper clay cap, indicating
temperature less than 140°C is clearly correlating to Figure 7 and extends up to 700-800 m depth, which are cased off in most of the wells.

4. WELL TEST ANALYSIS

4.1 Introduction

In an injection step test the injection to a well is increased or decreased in steps while the pressure is monitored at a fixed location, fixed depth. The Welltester software (Martinson, 2017; Júlíusson et al., 2008; Horne 1995), which was developed at Iceland GeoSurvey for welltest analysis, is used to simulate pressure response data, measured during a step test (Figure 8) as a function of time. Welltester deduces reservoir and well parameters by iterating, starting with guessed parameters and getting closer and closer to the measured pressure.

Axelsson and Steingrímsson, (2012) briefly explained about well test which is a fluid flow test (Figure 8) conducted in geothermal wells to obtain parameters of the reservoir as well as the well. Well tests are done at the completion of a well, possibly leading to a decision of continuing drilling, but also after a period of production, to see whether and how much the reservoir properties have changed.

A well test may be analysed using Theis model and its variants which assumes a model of homogeneous, isotropic and horizontal permeable layer of constant thickness, confined aquifer, and two dimensional and horizontal flow towards the producing well by fitting the pressure response of the model to the measured pressure response data (Axelsson, 2013). In a similar way the flow in an
injection well test is assumed to be horizontal from the well to the surroundings. The possible boundary conditions are illustrated in Figure 9 below.

![Figure 9: Pressure response of Theis model on a semi-logarithmic plot (linear pressure change vs. logarithmic time) demonstrating linear behaviour, which is the basis of the semi-logarithmic analysis method Rödvarsson and Whiterspoon (1989)](image)

4.2 Pressure Diffusion Equation

According to Haraldsdóttir (2017), pressure diffusion equation is a mathematical description of fluid flow in porous medium and used to calculate the pressure ($p$) in the reservoir at a certain distance ($r$) from the producing well producing at rate ($Q$) after a given time ($t$) as shown in Figure 10 below. It consists of three main physical principles; the law of conservation of mass, Darcy’s law and equation of state of the fluid. There are several assumptions to develop the pressure diffusion equation.

![Figure 10: Radial flow in a cylinder around a wellbore Haraldsdóttir, (2017)](image)

The law of conservation of mass is based on continuity equation of fluid flow in a porous medium which can be expressed as:

$\frac{\partial \left( \rho Q \right)}{\partial r} = 2\pi r \frac{\partial \left( \rho \phi h \right)}{\partial t}$  \hspace{1cm} (3)

where

- $Q$ = Fluid flow (mass introduced (source) or mass removed (sink)) (L/s)
- $\rho$ = Density of the fluid (kg/m$^3$);
- $\phi$ = Porosity of the medium (%);
- $t$ = Time(s);

The momentum equation (Darcy’s law) expresses the fact that the volumetric rate of flow at any point in a uniform porous medium is proportional to the gradient of potential in the direction of flow at that point.

$Q = \frac{2\pi k h \frac{\partial p}{\partial r}}{\mu}$  \hspace{1cm} (4)

where

- $P$ = pressure (bar)
The third principle is the equation of state of the fluid (fluid compressibility at constant temperature). The compressibility of a substance is the change in volume per unit volume per unit change in pressure. In a reservoir which consists of rock ($C_r$) and pore space occupied by oil, water, and gas ($C_w$), the total compressibility ($C_t$) is defined as follows:

$$C_t = C_w + (1-\varphi)C_r$$

Finally, the pressure diffusion equation will be deduced to a one dimensional second order partial differential Equation (6) which describes isothermal flow of a fluid in porous media, i.e. how the pressure (p) diffuses radially through the reservoir as a function of the distance ($r$) from the well and the time ($t$) since the start of production.

$$\frac{1}{r} \frac{\partial}{\partial r} \left( r \frac{\partial p}{\partial r} \right) = \frac{\mu \phi}{k} \frac{\partial p}{\partial t} = \frac{S}{T} \frac{\partial^2 p}{\partial r^2} + \frac{1}{r} \frac{\partial p}{\partial r} = \frac{\mu \phi}{k} \frac{\partial p}{\partial t} = \frac{S}{T} \frac{\partial^2 p}{\partial r^2}$$

Axelsson and Steingrímsson(2012) explained in detail the parameters of a well and reservoir to be estimated from well-logs during step-rate well-test and its purpose, which is conducted at the end of drilling a well. Step-rate tests are done to obtain a first estimate of the possible production capacity of a well and to estimate its production characteristics. In the case of high-temperature wells this estimate is only indirect since it’s not performed at high-temperature, production conditions. Step-rate well-testing usually lasts from several hours to a few days.

### 4.3 Injection step test analysis of well LA-4

Three injection rates are performed during the multi rate injection test analysis. The injection rates are 6.6 L/s, 17.2 L/s and 27.5 L/s. The results from the simulations of Welltester are shown in Table2 for the first step and the fall off step. The coefficient of variation (CV) for the two steps is within reasonable range except for the storativity ($S$) in the first step and radius of investigation ($r_e$) from the well and the time ($t$) since the start of production.

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Step 1 Value</th>
<th>Step 1 CV (%)</th>
<th>Fall off Step Value</th>
<th>Fall off Step CV (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Transmissivity, T (m³/(Pa.s))</td>
<td>4.8*10⁻⁹</td>
<td>2.8</td>
<td>2.1*10⁻⁹</td>
<td>1.1</td>
</tr>
<tr>
<td>Storativity, S (m³/(m²Pa))</td>
<td>1.8*10⁻⁸</td>
<td>26</td>
<td>3.2*10⁻⁸</td>
<td>4.6</td>
</tr>
<tr>
<td>Radius of Investigation, $r_e$ (m)</td>
<td>100</td>
<td>30</td>
<td>250</td>
<td>9</td>
</tr>
<tr>
<td>Skin factor, $s$</td>
<td>-3.0</td>
<td>-0.7</td>
<td>-0.7</td>
<td>-0.7</td>
</tr>
<tr>
<td>Wellbore storage, C (m³/Pa)</td>
<td>4.3*10⁻⁹</td>
<td>2.4</td>
<td>4.2*10⁻⁹</td>
<td>- 0.7</td>
</tr>
<tr>
<td>Reservoir Thickness, h (m)</td>
<td>260</td>
<td>450</td>
<td>450</td>
<td>2.2</td>
</tr>
<tr>
<td>Injectivity Index, II ((L/s)/bar)</td>
<td>1.5</td>
<td>2.2</td>
<td>2.2</td>
<td>2.2</td>
</tr>
</tbody>
</table>

### 4.4 Injection step test analysis of well LA-10D

This is the second directional well drilled in Aluto Langano geothermal project located in the central part of the main Ethiopian rift system, which is about 200km away from Addis Ababa. It is the 10th deep well drilled in Aluto volcanic complex completed drilling in October 2015. The well is drilled to the maximum geological depth of 1951m in the direction of N43°W with an inclination of 27.75° and a maximum temperature of 310°C at the bottom of well. The directional drilling kick of point (KOP) is 450m. The PTS tool was set at 551m depth and the water injection rate increased from 0 l/s to 14.9 L/s. Then, the injection rate was increased to 17, 19.34 and 21.73 L/s steps and the pressure in the well recorded. The injection period of each step was one hour (Figure 11).
The reservoir parameters for all steps are summarized in Table 3 and Step three was selected as the best model, having lower CV values than other steps.

### TABLE 3: Summary of Reservoir parameters estimated using nonlinear regression model for well LA-10D

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Step 1</th>
<th>CV %</th>
<th>Step 2</th>
<th>CV %</th>
<th>Step 3</th>
<th>CV %</th>
</tr>
</thead>
<tbody>
<tr>
<td>Transmissivity, T (m³/(Pa.s))</td>
<td>1.26×10⁻⁹</td>
<td>0.71</td>
<td>1.99×10⁻⁸</td>
<td>0.71</td>
<td>1.32×10⁻⁸</td>
<td>0.41</td>
</tr>
<tr>
<td>Storativity, S (m³/(m²Pa))</td>
<td>5.1×10⁻⁹</td>
<td>2.17</td>
<td>3×10⁻⁸</td>
<td>3.45</td>
<td>5.7×10⁻⁹</td>
<td>2.13</td>
</tr>
<tr>
<td>Radius of investigation, r (m)</td>
<td>122</td>
<td>5.61</td>
<td>62</td>
<td>2.45</td>
<td>119</td>
<td>4.81</td>
</tr>
<tr>
<td>Skin factor, s</td>
<td>0.07</td>
<td>-</td>
<td>3.08</td>
<td>-</td>
<td>-0.26</td>
<td>-</td>
</tr>
<tr>
<td>Wellbore storage, C (m³/Pa)</td>
<td>3.08×10⁻⁶</td>
<td>0.68</td>
<td>4.13×10⁻⁶</td>
<td>0.24</td>
<td>4.31×10⁻⁶</td>
<td>0.34</td>
</tr>
<tr>
<td>Reservoir thickness, h (m)</td>
<td>152</td>
<td>-</td>
<td>87</td>
<td>-</td>
<td>168</td>
<td>-</td>
</tr>
<tr>
<td>Injectivity index, II ((L/s)/bar)</td>
<td>1.41</td>
<td>-</td>
<td>1.4</td>
<td>-</td>
<td>1.63</td>
<td>-</td>
</tr>
</tbody>
</table>

Injection well test analyses done for wells LA-4 and LA-10D gave the best result for fall off step and step 3 respectively. The model assumes the reservoir as homogenous reservoir and constant boundary pressure. The skin factor values for the selected model step is negative, which indicates that the nearest surroundings of the wells have higher permeability than the surrounding reservoir and are in good connection with the reservoir. Both wells LA-4 and LA-10D are characterized to be in liquid dominated reservoir.

### TABLE 4: Comparing estimated reservoir parameter results of LA-4 and LA-10D of best model

<table>
<thead>
<tr>
<th>Parameter</th>
<th>LA-4</th>
<th>LA-10D</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Fall off step</td>
<td>CV %</td>
</tr>
<tr>
<td>Transmissivity, T (m³/(Pa.s))</td>
<td>2.1×10⁻⁸</td>
<td>0.37</td>
</tr>
<tr>
<td>Storativity, S (m³/(m²Pa))</td>
<td>3.2×10⁻⁸</td>
<td>1.16</td>
</tr>
<tr>
<td>Skin factor, s</td>
<td>-0.65</td>
<td>-</td>
</tr>
<tr>
<td>Wellbore storage, C (m³/Pa)</td>
<td>4.2×10⁻⁶</td>
<td>0.45</td>
</tr>
<tr>
<td>Reservoir thickness, h (m)</td>
<td>450</td>
<td>168</td>
</tr>
<tr>
<td>Injectivity index, II ((L/s)/bar)</td>
<td>2.2</td>
<td>1.63</td>
</tr>
<tr>
<td>Porosity, φ</td>
<td>0.1</td>
<td>0.1</td>
</tr>
</tbody>
</table>

### 5. CONCLUSION

Aluto Langano geothermal field is classified as high temperature geothermal field as the temperature is greater than 200°C at a depth of 1000 m, which was confirmed in this project. The main conclusions can be summarised as follows:

- Results from LA-9D shows an estimated formation temperature ranging from 300 – 311°C in the well bottom. The pivot point of LA-9D is located around 1100m-1250m depth with pressure ranging from 85-95 bar.
- Whereas the formation temperature of LA-10D is 300-305 °C while the pivot point is located around 1200-1400m depth with pressure ranging from 85-100 bar.
- The formation temperature reaches over 300°C in LA-6 and 230°C in LA-4. LA-4was drilled in the eastern and colder part of the geothermal field compared to wells LA-6, LA-9D and LA-10D, which were drilled in the warmest part of the field. This confirms results already shown in previous reports about the Aluto Langano geothermal field (ELC, 2016, WestJEC, 2016).
- The Welltest software used in this project to manipulate and analyse the injection step tests was helpful in estimating the reservoir parameters of the wells. The results from the analysis in wells LA-4 and LA-10D were fairly good, especially with respect to the data quality, which was scarce and limited. After manually fixing the data and using interpolation, the best simulations were for dual porosity reservoir, constant pressure boundary and constant skin.
The transmissivity and storativity of wells LA-4 and LA-10D were found to be in the order of magnitude of $10^{-8}$, which are similar or lower than in Icelandic geothermal wells, where they are in the order of magnitude $10^{-8}$.

Aluto Langano is classified as a high temperature geothermal field but the injectivity index is rather low, in the range of 1.0 to 2.4 (L/s)/bar which indicates that the wells as being of relatively moderate permeability. Low values of transmissivity indicate low permeability of the system. The field is therefore not necessarily a good producer in spite of the high temperature. This is in agreement with results of WestJEC (2016). The reservoir thickness for well LA-4 is 450m while for LA-10D is around 170m.

REFERENCES


