

San Juan single-well seismic data analysis and modeling study

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Summary

We analyze single-well seismic data from the San Juan Basin in Northwest New Mexico. The consistently observable events are tube-waves: direct, reflected and multiple tube-waves can be explained by the formation properties and survey geometry except for an anomalous zone with low velocity, high amplitude and horizontal polarization.

To aid the data analysis, forward modeling using a variable-grid finite-difference parallel code is performed. The numerical result confirms the identified events in the field observations.

Introduction

Sponsored by the U.S. Dept. of Energy, a large study on fracture quantification in naturally fractured gas reservoirs has been undertaken (Majer et al., 2001). As part of the project, field data at different scales (logs, single-well, VSP, and surface seismic) were collected in the San Juan Basin in Northwest New Mexico (Majer et al., 2002). We are conducting the single-well seismic data analysis and modeling study.

In single-well field data, tube-waves are the consistently observable events, though weak direct S-waves and tube-wave to body-wave conversions at interfaces can be seen on some sensors. Here we present the analysis on (1) the generally visible events; and (2) an anomalous zone with low velocity and high amplitude. We also present a numerical modeling result that confirms the data analysis.

Survey information

For the San Juan single-well experiments, two surveys were attempted: one with the orbital vibrator source and 3-C (3-component) geophone receiver; the other with a high frequency piezoelectric source and hydrophone receiver. Figure 1 shows the equipment geometry. We are focusing on the orbital vibrator data analysis since only limited piezoelectric data were acquired.

For the orbital vibrator survey, the source depth ranges from 4245 to 4900 ft at 5 ft intervals. There are four 3-C (H1, H2 and V) receivers per shot at offsets 87, 97, 107, and 117 ft. The orbital source with clockwise and counter-clockwise rotations at each location is decomposed into two orthogonally horizontal components (H1 and H2); this was performed at LBNL. Therefore, six components are recorded for each source depth. Shot gathers have one source depth for four 3-C receivers. Common offset

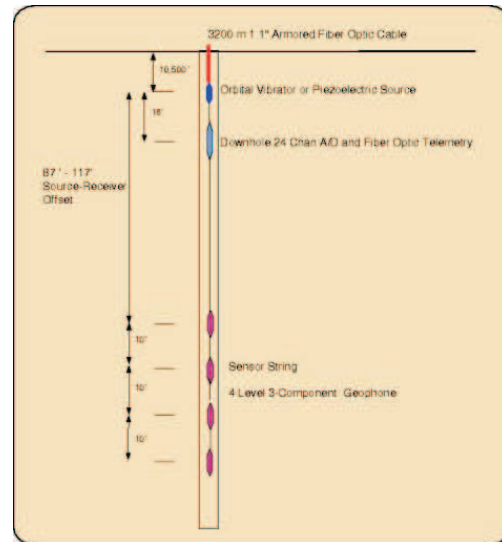


Fig. 1: Single-well equipment geometry.

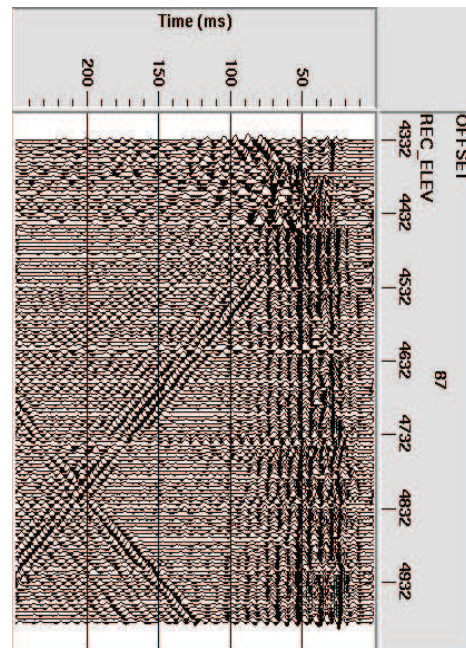


Fig. 2: One component (source H2, receiver V) of the common offset gather of channel 1 (offset 87 ft) with AGC.

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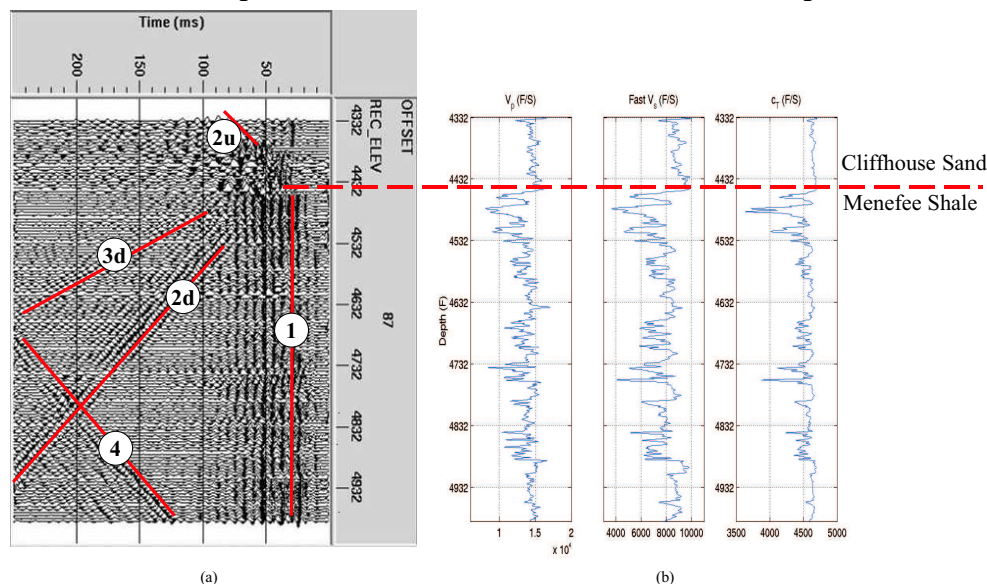


Fig. 3: (a) Generally visible events in the field data; (b) Sonic logs (P- and fast S-waves) and the estimated low frequency tube-wave velocity from logs. Velocities are used to predict the arrival times to identify the events in the data: 1 - direct tube-wave; 2u/2d - up/down-going interface reflections; 3d - down-going multiple reflections; 4 - reflected tube-wave from the well plug.

gathers have a single source-receiver pair for all source depths. We use common offset gathers for data analysis since they have the whole depth coverage.

Field data analysis

Figure 2 shows one component (source H2, receiver V) of the common offset gather of channel1 (offset 87 ft). Direct and reflected waves are generally visible. A low velocity zone between depth 4387-4442 ft is observed. We now present analysis of these.

Events analysis

Logs and survey geometry can help us to identify the events observed in the field data (Figure 3a).

Figure 3b shows sonic logs (P- and fast S-waves) and the estimated low frequency (<1000Hz) tube-wave velocity (White, 1983). Using these velocities, we predict the arrival times of the direct P-, S- and tube-waves which indicate that the consistently observable direct arrivals (event 1) in the field data are tube-waves, though weak S-waves are visible on some sensors. Primary tube-wave reflections, events 2u/2d and 4, are from the C/M interface (Cliffhouse sand/Menefee shale interface) and the well plug, respectively. Event 3d is due to multiple reflections between the tool and the C/M interface.

Figure 4 schematically illustrates the wave propagation paths (Figure 4a) and associated arrivals on the seismogram (Figure 4b). Except the direct tube-waves (1) and the well plug reflections (4), the C/M interface reflections and multiples are affected by the tool position. When the

tool is above the interface, there are up-going primary reflections from the interface (2u) and multiple reflections between the tool and the interface (3u). When the tool is across the interface (the source is above and the receiver is below the interface), neither primary nor multiple reflections will be observed (dash lines in Figure 4b). When the tool is below the interface, there are down-going interface reflections (2d) and multiple reflections (3d).

For a common offset gather, slopes of events on the seismogram are their apparent velocities. Using the ray paths for source/receiver at two depths, we get the formulation to calculate the true velocities of the events in Figure 4b:

$$\begin{aligned} c_1 &= \text{S-R offset} / \Delta t \\ c_{2u} &= c_{2d} = c_4 = 2 * \Delta h / \Delta t = 2 * c' \\ c_{3u} &= c_{3d} = 4 * \Delta h / \Delta t = 4 * c', \end{aligned} \quad (1)$$

where $c' = \Delta h / \Delta t$, called apparent velocity, is the slope of the event on the seismogram.

Figure 5 shows the apparent and true velocities of the identified events in the field data. We can see these events have almost the same true velocities that confirm they are all tube-waves.

Anomalous zone analysis

In the field seismogram (Figure 2), for receiver depths between 4387-4442 ft, the observed direct wave has a very low velocity. True amplitude display (Figure 6) shows the event has a very high amplitude. The following is the analysis of this anomalous zone.

In figure 3, we can see the anomalous zone is right above

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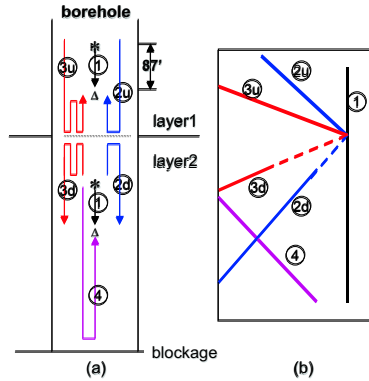


Fig. 4: Schematic diagram of tube-waves in the single-well survey: (a) Wave paths; (b) Arrivals on the seismogram.

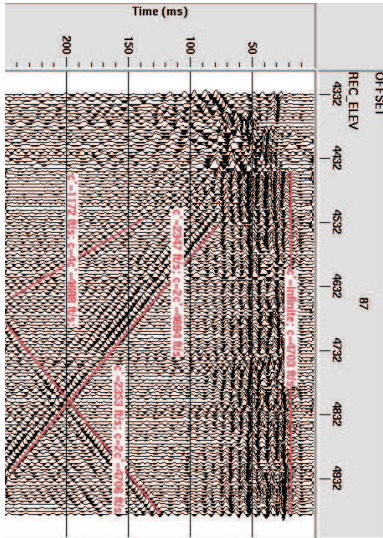


Fig. 5: Apparent velocity (c') and true velocity (c) of the direct, reflected, and multiple tube-waves in the field data with AGC

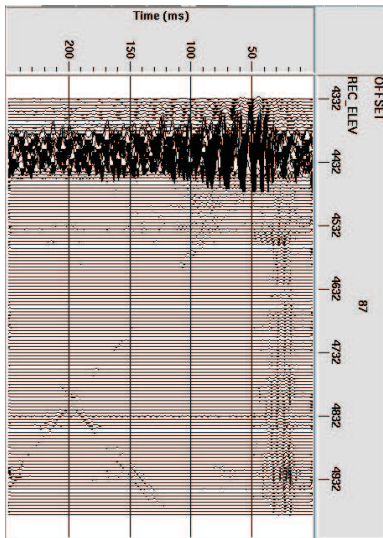


Fig. 6: True amplitude display of the data: the anomalous zone (4387-4442 ft) with low velocity and high amplitude.

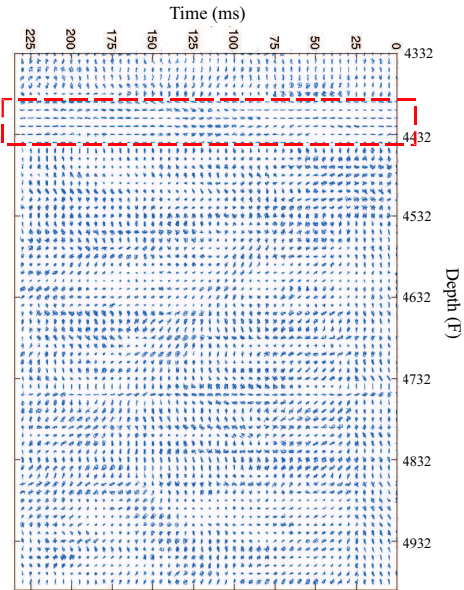


Fig. 7: Polarization diagram for receiver H2 and V of the source H2. The anomalous zone (in the red box) has horizontal polarization but the rest has vertical polarization.

the Menefee shale and is in the Cliffhouse sand. The predicted tube-wave velocity from logs (the 3rd panel in figure 3b) for this zone is around 4700 ft/s which is much higher than the velocity of the observed direct wave on the field seismogram (2900 ft/s). Therefore formation properties obtained from logs do not explain the anomaly.

The particle motion study (Figure 7) shows the anomalous event is horizontally polarized, but the normal tube-wave is vertically polarized. The horizontal polarization is consistent with flexural-waves, but the dispersion relation of flexural-waves (Paillet and Cheng, 1991) indicates that the velocity of the flexural-wave at the single-well frequency (300-400 Hz) can not be as low as that of the anomalous event in the field observations.

Comparing the polarization of the four common-offset gathers (offsets 87, 97, 107 and 117 ft) by fixing the source and receiver depth, respectively, we find the anomaly only happens when the source is below 4300 ft and the receiver is above 4442 ft. This indicates that it is not caused by either the source or receiver, but is related to some formation properties not explained by open-hole logs. Out of plane vertical gas-filled fractures can be one possibility, but more study on the amplitude of gas-filled fracture reflections are needed.

Another interesting finding is that the anomalous zone shows direct correspondence with the highest permeability zone interpreted from open hole logs. However, a mechanism which could explain the relationship is not yet determined.

Modeling study

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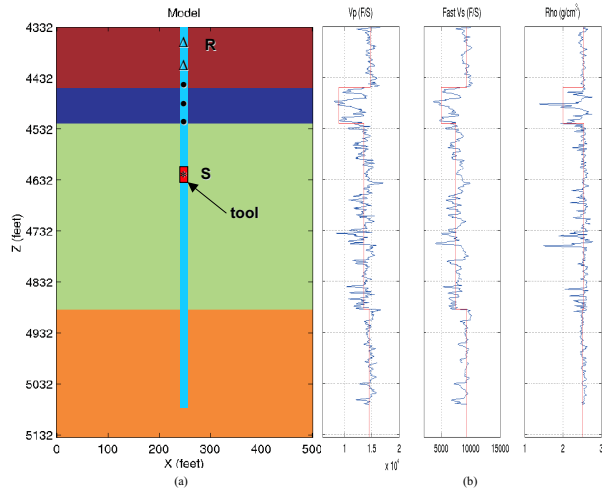


Fig. 8: (a) Model; (b) Blocked logs.

To aid the data analysis, modeling study using a 2-D variable-grid finite-difference parallel code is performed. Here we present one result to confirm the identified events in the field data.

Figure 8a shows the layered model with formation properties from the logs (figure 8b). The water-filled borehole is blocked at 5082 ft. A x direction dipole source, at 4622 ft, is used to excite the model. A small rectangular steel box is placed around the source to represent the tool. Unlike the field survey with four receivers per shot, in the simulation, we put receivers on all grids along the borehole for one shot. The synthetic common shot gather with the whole depth coverage is used to assist the field data analysis since it is computationally costly to generate the same synthetic common offset gather as the field data. A variable grid mesh, which efficiently handles the small borehole in a large model, is used for the modeling.

The synthetic common shot gather, figure 9, shows tube-waves dominate the wavefield though S-waves are visible in some traces. Direct tube-waves (T), C/M interface and well plug reflections (TT_i and TT_p), and multiple reflections between the tool and C/M interface (TT_iT_i) observed in the synthetic seismogram confirms the identified events in the field data.

Conclusions

We have conducted single-well acquisition, data analysis and modeling. The data analysis demonstrates that the consistently observable events in the field data are tube-waves. Except an anomalous zone with the low velocity and high amplitude, the direct, reflected and multiple tube-waves can be explained by the formation properties from logs and the survey geometry. Numerical modeling confirms our data analysis.

Future work will involves analyzing the events at the depths with fracture indications from FMI (Formation

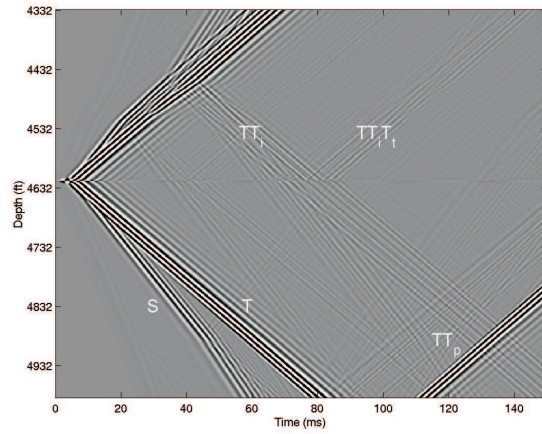


Fig. 9: Synthetic common-shot gather: *S* - direct S-wave; *T* - direct tube-wave; TT_i - C/M interface reflected tube-wave; TT_p - well plug reflected tube-wave; TT_iT_i - multiple tube-wave reflections between the tool and the interface.

Micro Imager).

Acknowledgments

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