

PROPOSAL FOR RESERVOIR ENGINEERING STUDIES
IN THE STATE OF ALASKA

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Alaska has a significant geothermal potential. While other sources of energy such as petroleum and coal are in abundance, there has been a definite move towards geothermal exploitation. The State has recognized the opportunity cost of petroleum as a source of materials and has expressed interest in the development of geothermal energy as a desirable and alternative resource.

More than 11 million acres ¹ have been identified as potential geothermal reservoirs capable of producing electric power as well as direct heating. Reservoirs of the latter type are found in the interior of the state. Considering the winter temperatures of these regions (at times dipping to -60°F) direct utilization is attractive.

A comprehensive reservoir engineering proposal is presented to better assess the extent and potential of the geothermal areas in Alaska. The purpose of this paper is to acquaint the participants of the Stanford Geothermal Workshop with the enormous potential, as yet untapped, of the State of Alaska.

Introduction

Figure 1 is a map of Alaska with a number of potential geothermal sites. Of these reservoirs, Chena Hot Springs and Pilgrim Hot Springs have had considerable geological and geophysical work. ^{2,3} At Pilgrim Springs, two small wells were drilled at 200 feet and they encountered hot water at 178°F. No deeper drilling was attempted since the available equipment was incapable for such a task. The flow rate was measured at 32,000 lb/hr. At Chena Hot Springs and at a depth of only 18 feet a temperature of 138°F was recorded.

Two other exploration projects are to begin in calendar 1981. Personnel from the Geophysical Institute and the Petroleum Engineering Program at the University of Alaska, Fairbanks will cooperate in the "Geophysical Exploration for Geothermal Energy at Manley Hot Springs" and in the "Investigation of Radiogenic Heated Aquifers in the Lower Susitna Basin".

Potential sites in Alaska are: Pilgrim Hot Springs, Kotzebue, west side of Mount Drum (Klawasi), Willow, Chena Hot Springs, Circle Hot Springs, Manley Hot Springs, Homer Hot Springs, Clear Creek Hot Springs, Central Baranof Island, Tenakee Hot Springs, Northern part of Unalaska Island, Unnak Islands, Emmons Caldera and Northeastern Atka Island. (Reference 1)

As it can be seen, the developable sites in the state span the entire area.

Many of these areas occur in sedimentary basins, containing ancient volcanic formations. Often, geothermal formations may be associated with hydrocarbon deposits. Fractures, which are characteristic of other major geothermal formations, also penetrate the Alaskan reservoirs.

RESERVOIR ENGINEERING WORK

The University of Alaska will purchase Amerada bombs in order to facilitate a comprehensive well testing program. Following intensive geophysical exploration it is anticipated that an aggressive drilling program will ensue. Drawdown and buildup well tests will be done in the first well. If they are encouraging then a second and third well will be drilled preferably forming a right angle with the first well. One direction should follow identified faults while the other will be perpendicular. Interference well testing will allow the estimation of the directional permeabilities and the identification of the principle axes of permeability.

Subsequent drilling will take into account the findings of the well testing. Well logging during the drilling phase will supply the final formation evaluation which along with the results of well testing will define the economic attractiveness of each project.

While reservoir engineering work in the state is still in its infancy, it is expected to grow rapidly. The comprehensive geophysical and geological work done thus far have prepared a highly favorable ground for the reservoir engineer.

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3. Turner, Donald L., Wescott, Eugene and Kienle, Juergen,: "A Geological and Geophysical Study of the Chena Hot Springs Geothermal Area, Alaska", Geophysical Institute Preliminary Report, University of Alaska, Fairbanks, 1980.

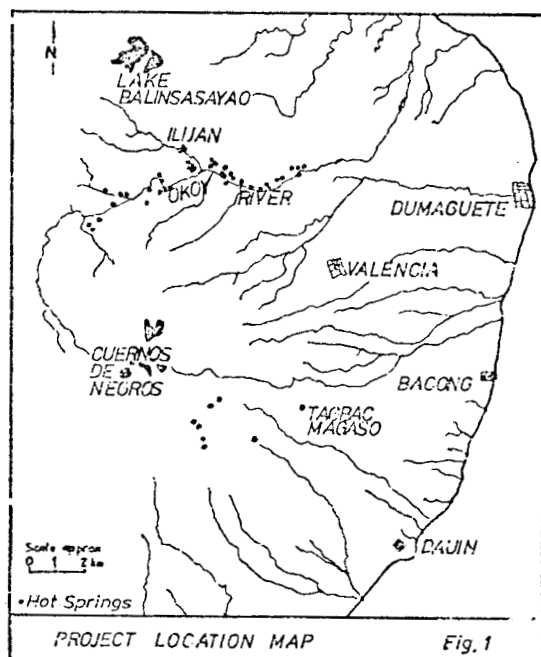
Figure 1. Alaska map with potentially developable sites.

HYDROLOGY AND MODEL OF THE OKOY GEOTHERMAL FIELD,
NEGROS ORIENTAL, REPUBLIC OF THE PHILIPPINES

E.W. Smith, Geothermal Engineer, New Zealand

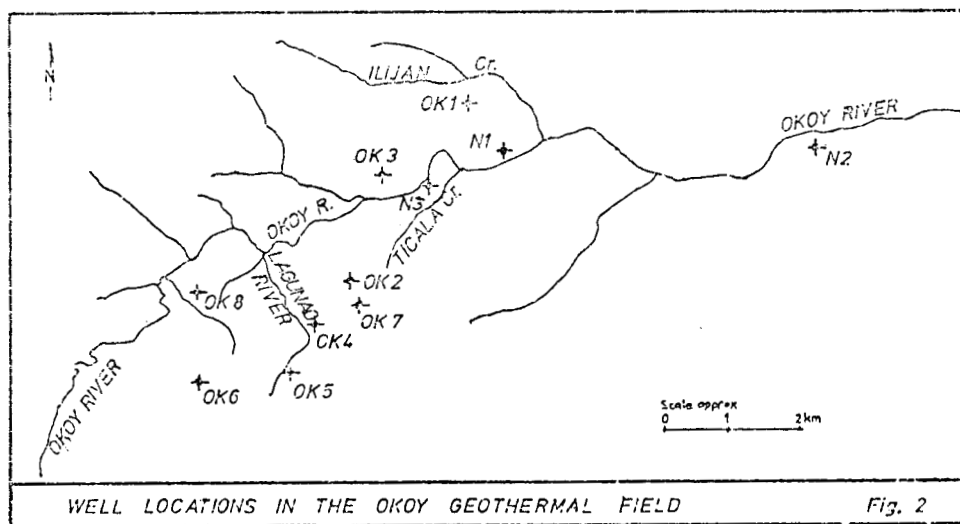
INTRODUCTION

Ward (1980) described the exploration of the Okoy geothermal field. Resistivity surveying using Schlumberger traverses has covered an area of approximately 800 km² in the southern part of the Island of Negros. Hot springs and other thermal manifestations (Fig. 1) occur in the Okoy valley, Valencia and at Tabac Magaso, Dauin.



Initial shallow exploratory drilling indicated a possible sub-surface flow of hot water in the Okoy valley. Further deep exploratory drilling to the West has located two high temperature reservoirs.

Figure 2 shows the well locations within the Okoy geothermal field.



HYDROLOGY

Due to the complex nature of the Okoy geothermal field, wells drilled into the hydrothermal system generally balance well and reservoir pressures with internal flows. These internal flows cause difficulty in interpreting the downhole temperature and pressure measurements.

Most of the wells within the field penetrate a single phase reservoir and some have multiple permeable zones. Wells with multiple permeable zones generally have one zone with larger permeability and at this point the measured pressures equal the reservoir pressure. This point is known as the pressure control point. Elsewhere in the well the measured pressures do not equal the reservoir pressure.

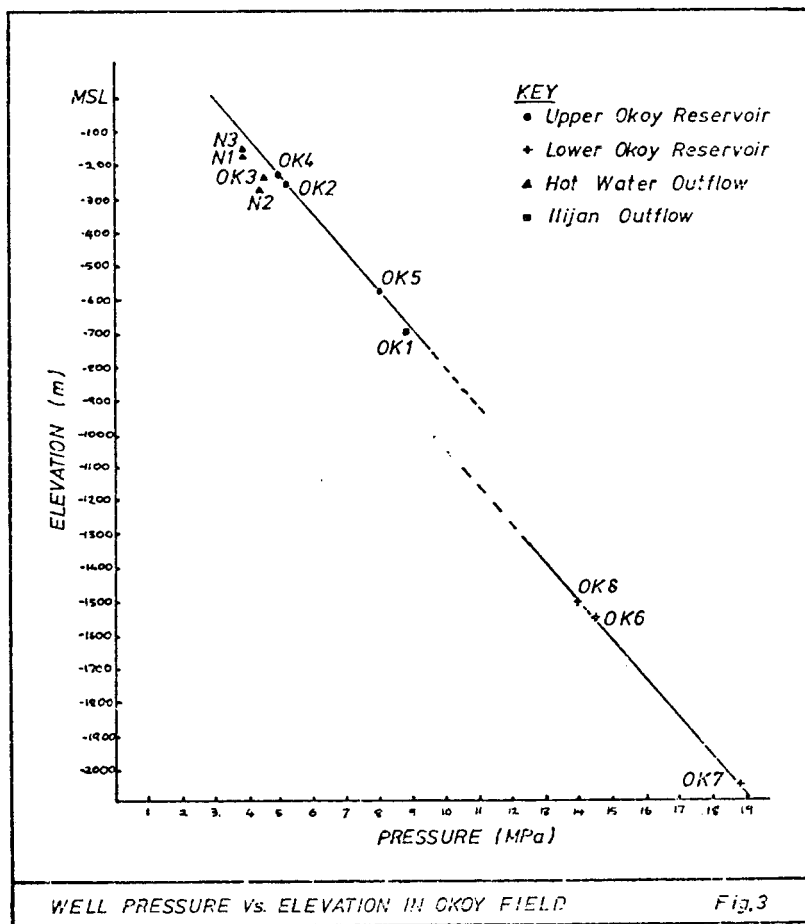
After examination of the well measurement data the pressure control point for each well was identified and listed in Table 1.

The pressure survey data was examined and rejected where the surveys showed transient effects, strong internal flows or obvious measurement errors. The remaining surveys were used to calculate the average pressures at the pressure control point for each well and listed in Table 1.

Table 1: Elevation - Pressure Data

Well	Pressure Control Point m MSL*	Average Pressure kPa
N1	- 169	3870
N2	- 268	4360
N3	- 152	3860
OK1	- 687	8810
OK2	- 246	5170
OK3	- 232	4500
OK4	- 223	4930
OK5	- 568	8020
OK6	- 1544	14540
OK7	- 2040	18810
OK8	- 1496	14000

The elevation - pressure data of Table 1 is graphed in Figure 3.



* MSL = mean sea level

Examination of Figure 3 and a knowledge of the physical geography, geology and chemistry of the Okoy field suggests that the data can be grouped into four categories:-

- (a) Upper Okoy reservoir (OK2, OK4, OK5)
- (b) Lower Okoy reservoir (OK6, OK7, OK8)
- (c) Hot water outflow (N3, OK3, N1, N2)
- (d) Ilijan outflow (OK1)

UPPER OKOY RESERVOIR

The upper Okoy reservoir is a single phase system existing between +300m and -1000m MSL consisting of neutral chloride water with chloride concentrations of approximately 3600ppm. Deep drilling has encountered temperatures of up to 310°C and no re-charge system has yet been identified.

Some small two phase steam dominated zones overlies the reservoir.

Least square linear regression analysis gives an 'elevation - pressure relationship' for this category (a) data of Table 1 as $P = 2960 - 8.91 z$. Where P is pressure (kPa) and z is elevation (m) relative to mean sea level. The co-efficient of determination for the analysis is 1.000 indicating an excellent fit.

The implications of the 'elevation-pressure relationship' are:-

- (i) The piezometric water level of the reservoir is +330m MSL.
- (ii) All Chloride springs which derive fluid from the reservoir will occur at elevations less than +330m MSL.
- (iii) Surface thermal manifestations at elevations greater than +330m MSL which derive fluid from the reservoir are likely to be steam heated or discharge free steam.
- (iv) The pressure gradient of 8.91 kPa/m corresponds to water of approximately 160°C. The average reservoir temperature is greater than 160°C thus the pressure gradient is super-hydrostatic.

LOWER OKOY RESERVOIR

The lower Okoy reservoir is a neutral chloride water system at elevations below -1000m MSL. Deep drilling so far has encountered temperatures ranging from 250 to 303°C and no recharge or outflow systems have been indicated.

Least squares linear regression analysis gives an 'elevation-pressure relationship' for the category (b) data of Table 1 as $P = 970 - 8.75 Z$. The co-efficient of determination for the analysis is 0.999 indicating a reasonable fit.

The implications of the 'elevation-pressure relationship' are:-

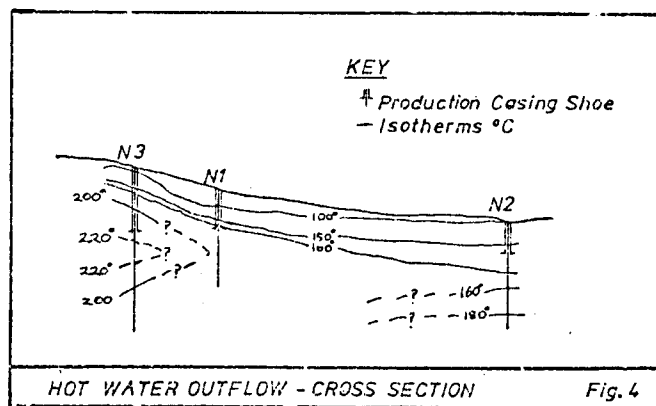
- (i) The piezometric water level of the reservoir is +110m MSL.
- (ii) All chloride springs which derive fluid from the reservoir will occur at elevations less than +110m MSL.
- (iii) The pressure gradient of 8.75 kPa/m corresponds to water of approximately 175°C. The average reservoir temperature is greater than 175°C thus the pressure gradient is super-hydrostatic.
- (iv) At an elevation of -1000m MSL the pressure difference between the upper and lower reservoirs is 2150 kPa. If permeability exists between these reservoirs, within the known field area, fluid will flow from the upper to the lower reservoir.

HOT WATER OUTFLOW

The shallow wells N1, N2 and N3 penetrate a hot water flow which supplies the Okoy valley springs.

Geochemical evaluation (Barnett, 1977) of the Okoy valley wells and springs indicated a general westerly direction as the source of the fluid.

Figure 4 shows isotherms drawn through wells N3, N1 and N2 also suggests a source to the west.



The Okoy geothermal field cross sections (Smith, 1980 b) suggested wells N3, N1 and N2 are in an outflow from the Upper Okoy reservoir with OK3 in the outflow edge.

Comparison of the data from Table 1 for these category (c) wells with the derived 'elevation-pressure relationship' for the upper Okoy reservoir is given in Table 2. The pressure departures of Table 2 show a sequence of increasing pressure drop in the order: N3, OK3, N1, N2 indicating the flow direction.

Table 2: Pressure Comparison

<u>Well</u>	<u>Pressure Control</u> <u>Point m MSL</u>	<u>Average Pressure</u> <u>kPa</u>	<u>P=2960-8.91%</u> <u>kPa</u>	<u>Pressure</u> <u>Departure</u>
N3	- 152	3860	4314	- 450
OK3	- 232	4500	5030	- 530
N1	- 169	3870	4470	- 600
N2	- 268	4360	5350	- 990

ILIJAN OUTFLOW

The Okoy geothermal field cross sections (Smith, 1980 b) suggested OK1 is on the edge of a second hot water flow.

Comparison of the data from Table 1 for OK1 with the derived 'elevation-pressure relationship' for the upper Okoy reservoir is given in Table 3.

Table 3: OK1 Pressure Comparison

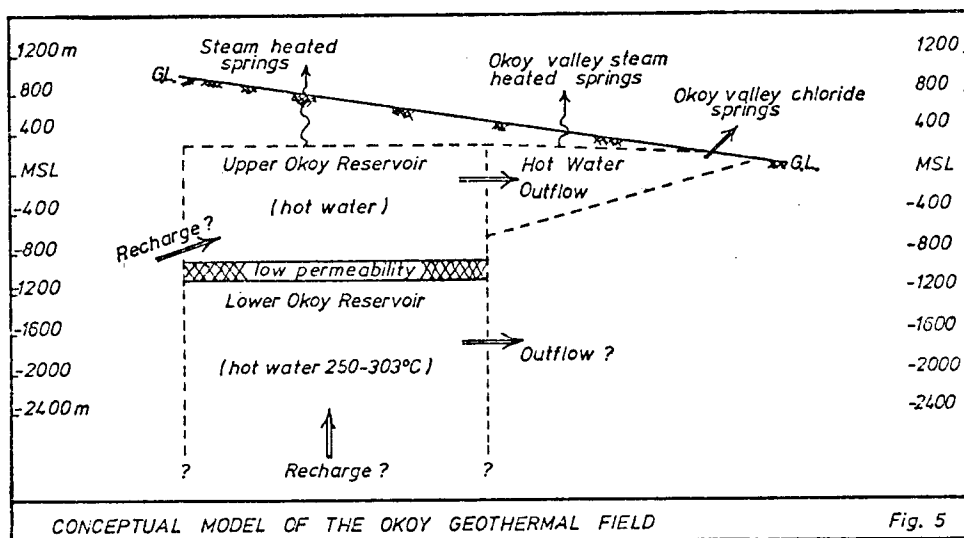
<u>Pressure Control</u> <u>Point m MSL</u>	<u>Average Pressure</u> <u>kPa</u>	<u>P=2960-8.91%</u> <u>kPa</u>	<u>Pressure</u> <u>Departure</u>
- 687	8810	9080	- 270

The pressure departure of Table 3 does not fit the sequence for the flow direction indicated by Table 2. It can be speculated that this second hot water flow may be from a geothermal system in the Balinsasayao area.

MODEL

A simple two dimensional conceptual model is given (Fig. 5) based on the descriptions of the upper and lower reservoirs and the hot water outflow.

Further exploration and delineation wells are being drilled in the Okoy geothermal field and it is hoped the information obtained will help clarify the complex nature of this field.



ACKNOWLEDGEMENTS

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